

Understanding the Spatiotemporal Structures in Atmosphere-Land Surface Exchange at the Jülich Observatory for Cloud Evolution (JOYCE)



Marke¹, T., S. Crewell¹, U. Löhnert, J. Schween, and U. Rascher²

¹ Institute of Geophysics and Meteorology, University of Cologne, ² Institute of Bio- and Geoscience, Research Center Jülich

1. Introduction to the Measurement Site

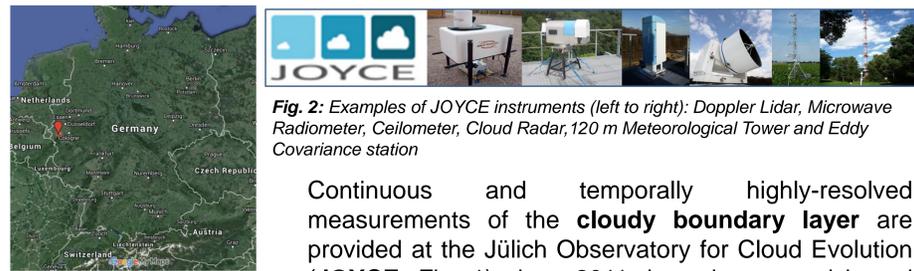


Fig. 1: Location of the JOYCE site in Jülich, Germany.

Fig. 2: Examples of JOYCE instruments (left to right): Doppler Lidar, Microwave Radiometer, Ceilometer, Cloud Radar, 120 m Meteorological Tower and Eddy Covariance station

Continuous and temporally highly-resolved measurements of the **cloudy boundary layer** are provided at the Jülich Observatory for Cloud Evolution (JOYCE, Fig. 1) since 2011, by using ground based passive and active remote sensing and in-situ instruments (Fig. 2).

2. Site Characterization with Ground Based Observations

- Macro-physical properties of boundary layer clouds are assessed with the **synergy** of a ceilometer and cloud radar.

- High monthly variability of the **total cloud cover** (Fig. 3)

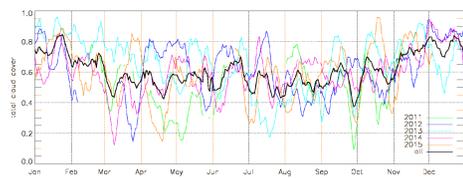


Fig. 3: Daily mean total cloud cover at JOYCE using a Ceilometer (10 day center moving average)

- Comparison of remote sensing, in-situ and model derived **wind direction** (Fig. 4)

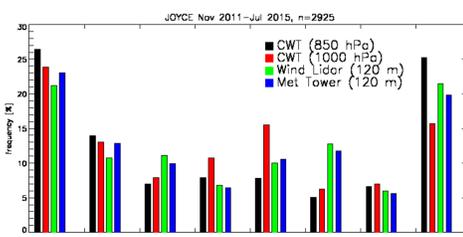


Fig. 4: Wind direction derived from Doppler Lidar and Meteorological Tower (120 m) and Circulation Weather Type Classification based on ERA-Interim 850 hPa / 1000 hPa.

- Influences of the local and synoptic scale can be identified

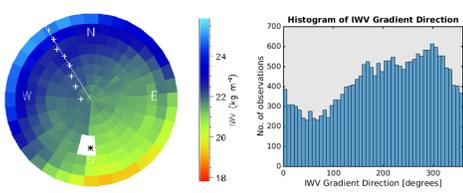


Fig. 5: Left: Air mass corrected IWV field and IWV gradient derived from one microwave radiometer hemispheric scan. Right: IWV gradient direction histogram.

- Azimuth scans using a microwave radiometer (Fig. 5) provide the spatial distribution and gradient of the **integrated water vapor (IWV)**.
- Link to exchange processes of the surface (Fig. 11)

3. A Scheme to Classify the Cloudy Boundary Layer

By reducing the observations of the boundary layer (BL) into a specific set of types helps understanding the **evolution of mixing processes** in the lowest part of the troposphere. Furthermore, the turbulence can be identified as cloud or surface driven (Harvey et al., 2013).

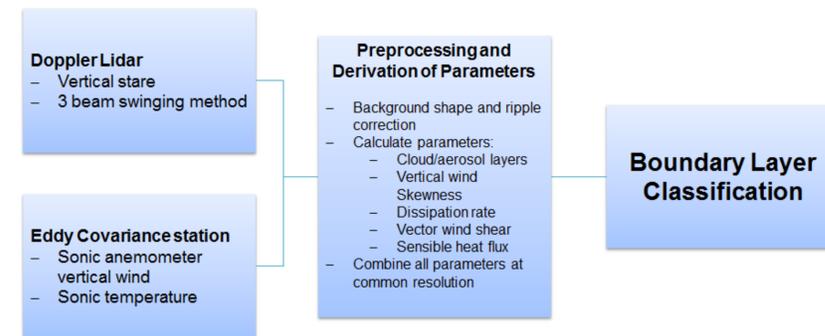


Fig. 6: Schematic representation of the boundary layer classification.

Preprocessing of Doppler Lidar Data

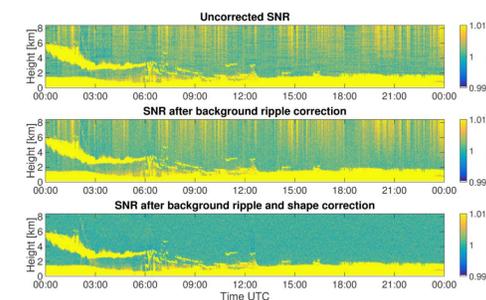


Fig. 7: Doppler Lidar SNR before and after correction.

Background shape correction (Manninen et al., 2015) and ripple correction (Vakkari et al., to submit)

- Homogeneous background
- Allows lower signal-to-noise ratio (SNR) threshold
- Bias in turbulent properties is reduced

Observations Used for the Classification

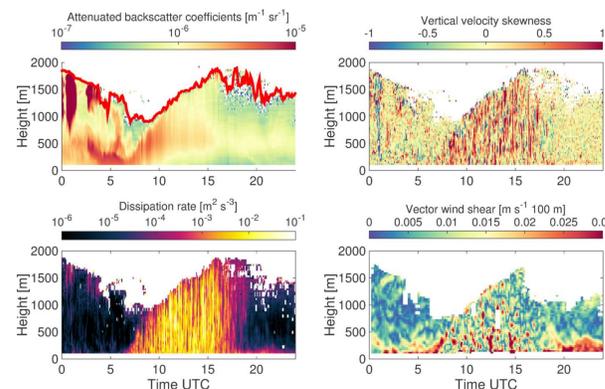


Fig. 8: Time series of the attenuated backscatter coefficient, vertical velocity skewness, dissipation rate and vector wind shear (clockwise from top left).

4. Boundary Layer Classification Results

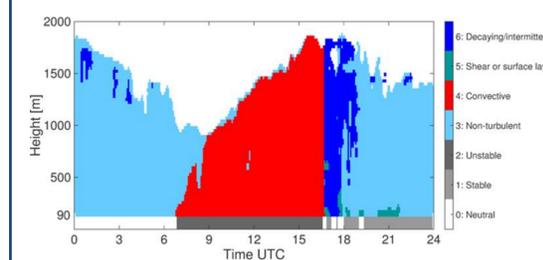


Fig. 10: Illustration of the bit-field, showing the boundary layer classification.

- Each pixel in the common resolution grid (temporal: 5 min, vertical: 30 m) is classified using a **bit-field**.

- Type decisions are based on combinations of **threshold values** (Table 1)

Parameter	Attenuated backscatter coefficient	Vertical velocity skewness	Dissipation rate	Vector wind shear	Sensible heat flux
Threshold	10^{-5}	0	$10^{-4} \text{ m}^2 \text{ s}^{-3}$ $10^{-3} \text{ m}^2 \text{ s}^{-3}$	0.02 m s^{-1} per 100 m	0 W m^{-2} $\pm 10 \text{ W m}^{-2}$

Table 1: Thresholds used for the bit-field.

Outlook: Operational use in the **Cloudnet** (Illingworth et al., 2007) framework.

5. Further Applications of the BL Classification

- HyPlant:** high-resolution airborne imaging spectrometer for vegetation monitoring (sun-induced and chlorophyll fluorescence, Fig. 11)

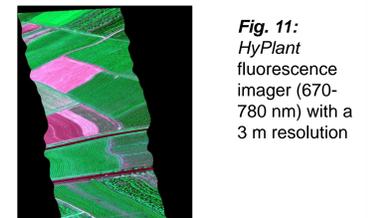


Fig. 11: HyPlant fluorescence imager (670-780 nm) with a 3 m resolution

- Link between IWV scans and surface patterns

- ICON** (ICOsaedral Nonhydrostatic): unified modeling system for global numerical weather prediction (NWP) and climate studies that performs as a large eddy simulation (LES) model (Fig. 12)
- Evaluate BL type parameterizations

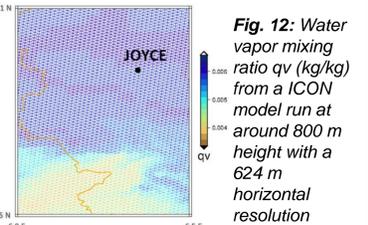


Fig. 12: Water vapor mixing ratio q_v (kg/kg) from a ICON model run at around 800 m height with a 624 m horizontal resolution

References:

Dipankar, A., B. Stevens, R. Heinze, C. Moseley, G. Zängl, M. Giorgetta, and S. Brdar, 2015: Large eddy simulation using the general circulation model ICON, *J. Adv. Model. Earth Syst.*, 7, 963-986.
 Harvey, N. J., R. J. Hogan, and H. F. Dacre, 2013: A method to diagnose boundary-layer type using Doppler lidar, *Q. J.R. Meteorol. Soc.*, 139, 1681-1693.
 Illingworth, A. J., and Coauthors, 2007: Cloudnet: Continuous evaluation of cloud profiles in seven operational models using ground-based observations. *Bull. Amer. Meteor. Soc.*, 88, 883-898.
 Manninen, A. J., E. J. O'Connor, V. Vakkari, and T. Petäjä, 2015: A generalised background correction algorithm for a Halo Doppler lidar and its application to data from Finland, *Atmos. Meas. Tech. Discuss.*, 8, 11139-11170.