Characteristics and Genesis Conditions of Polar Lows in between 2000-2012: Microwave satellites, Arctic System Reanalysis and Radiative Transfer Simulations



A. Radovan¹, S. Crewell¹, E. M. Knudsen¹, M. Mech¹, A. Rinke²

¹Institute of Geophysics and Meteorology, University of Cologne, ²Alfred Wegener Institute, Helmholtz Centre for Polar and Marine Research

TRANSREGIO TR 172 | LEIPZIG | BREMEN | KÖLN

UNIVERSITÄT LEIPZIG

Universität Bremen





Research Questions

RQ1. Is the Arctic System Reanalysis (ASR) able to represent polar lows (PLs) and their precipitation signature?

RQ2. Can we identify thresholds in environmental conditions or combinations of them that are required for PL formation?



Results

RQ1: Representation of PLs in AMSU-B and ASR

AMSU-B observations

- strong brightness temperature (BT) depression in precipitating ice cores
- BT difference to environment can reach more than 40 K

e (BT) cores	ASR integrated hydrometeor contents			AMSU- observations at 183.31+/-7	ASR using PAMTRA forward operator
	water vapor	cloud liquid	ice		
	IWV	LWC	ICE CON	183.31 \pm 7 GHz	183.31±7 GHz

Fig. 1: Distribution of January polar low cases (blue dots) between 2000-2012 using list of polar lows from Noer and Lien, 2010^[1]



Tools & Methods

<u>ASR v1</u> – Arctic System Reanalysis version 1 with 30 km spatial resolution and 29 vertical levels that has best estimate of atmospheric state including precipitation^[2] <u>Analyse 200 km around genesis point and time using:</u>

Conditions	Threshold
SST –T(500 hPa)	> 43 K ^[3]
SST – T(2m)	∼ 6 − 7 K ^[4]
Near surface wind speed	> 15 m/s ^[5]
RH (850 -950 hPa)	~ 82 % ^[4]
ΔMSLP _{mg}	≥ 1 hPa ^[6]
Lanse rate (LR) helow 850 hPa	Unstable ^[4]

AMSU-B simulations using PAMTRA

- general structure of the PL from ASR υτα is captured in the simulations
- BT signature difficult to see close to orography and sea ice due to emissivity change

ASR integrated values of PL

 general structure of the PL from ASR is visible in simulations

Possible reasons for the disagreement

- satellite has coarser resolution of the ASR (at nadir point doubled)
- parametrization of precipitation processes including assumptions of hydrometheor size and shape



Fig. 2: PL case on 7th, Jan, 2009 (top), 16th, Jan, 2009 (middle) and 8th, Jan, 2010 (bottom). Integrated water vapour (IWV) (first column), liquid water content (LWC) (second column), ice content (ICE CON) (third column); AMSU-B observations at 183.31±7 GHz channel (fourth column), PAMTRA simulations at 183.31±7 GHz channels (fifth column).

RQ2: Environmental conditions from ASR

	Genesis		Maturity
(a) -	SST - T(500 hPa)	(b)	SST - T(500 hPa)

Lapse rate (LR) below 850 hPa

Unstable

<u>Advanced Microwave Sounding Unit – B (AMSU-B) and</u> <u>Microwave Humidity Sounder (MHS</u>)

• coverage of the Arctic (\cong **10 times/day**) with 5 channels

2 window: 89 and 150 GHz (157 GHz MHS) 3 within strong water vapor line: 183.31 ± 1, 183.31 ±3, 183.31 ± 7 GHz (190 GHz MHS)

<u>PAMTRA</u> – Passive and Active Microwave Radiative TRAnsfer that connects ASR to AMSU-B and is able to simulate the 1-800 GHz frequency range using scattering





Fig. 3: Box-whisker representation (interquartile range in blue) of SST – T(500 hPa) (top) and lapse rate (LR) bellow 850 hPa (bottom) during genesis (left) and maturity stage (right) within a 220 km radius..

- higher amount of boundary layer rel. humidity during genesis stage (Fig. 4)
- more intense winds and lower MSLP at maturity stage
- boundary layer RH during PL days over region (Fig. 5) increased compared to January climatology



Fig. 5: RH profile for

- for the majority of the cases the SST – T(500 hPa) threshold of 43 K is reached (Fig 3).
- cases with stronger static
 stability show stronger and
 steeper lapse rates:
 → convection acts as driving mechanism





climatology (blue) and PL dates (red) over the whole region of investigation. Inserted figure is the difference between the two.

Fig. 4: Difference between genesis and maturity stage next variables:
MSLP difference (+), temperature at 2 m (*), near-surface wind speed
(NSWSmg ▲), SST (♦), and RH in the layer between 850 and 950 hPa (x).

Conclusions and next steps

- investigate the role of moisture intrusions or atmospheric rivers prior to a PL event
- analyze precipitation produced by PL
- **RQ1**: ASR transformed into observation space using forward simulator reproduces PL as detected by satellite measurements; validation technique difficult close to sea ice and orography
- **RQ2**: environmental conditions reveal the relative importance of thermal instability and convection for PL genesis; find the amount of precipitation brought by PL when making landfall

Acknowledgements

This work was supported by Transregional Collaborative Research Center TR172 "ArctiC Amplification: Climate Relevant Atmospheric and SurfaCe Processes, and Feedback Mechanisms (AC)³" funded by the German Research Foundation (DFG).

References

- [1] Noer and Lien 2010: Dates and Positions of Polar lows over the Nordic Seas between 2000 and 2010, NMI, no. 16/2010
- [2] ASR data, Polar Meteorology Group at ByrdPolar and Climate Research Center, the Ohio State University, available at https://rda.ucar.edu/datasets/ds631.1/
- [3] Noer, G. and M. Ovhed, 2003: Forecasting of polar lows in the Norwegian and Barents sea -proc. of the 9th meeting of the egs polar lows working grpup, Cambridge,
- [4] Terpstra et al., 2015: Forward and Reverse Shear Environments during Polar Low Genesis over the Northeast Atlantic, Monthly Weather Review, doi: 10.1175/MWR-D-15-0314.1 [5] Rasmussen, E. A. and J. Turner, 2003: Polar Lows. Cambridge University Press, Cambridge, UK

[6] Condron, A. and G. Bigg, 2006: Polar mesoscale cyclones in the NE atlantic: Comparing climatologies from era-40 and satellite imagery., Mon. Wea. Rev., 134, 1518–1533

POLAR2018, 18-23, June, 2018, Davos, Switzerland

Tue_69_AC-1_1820

Correspondence to: aradovan@uni-koeln.de