

Overlap statistics of shallow boundary layer clouds: Comparing ground-based observations with large-eddy simulations

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Boundary layer cloud overlap

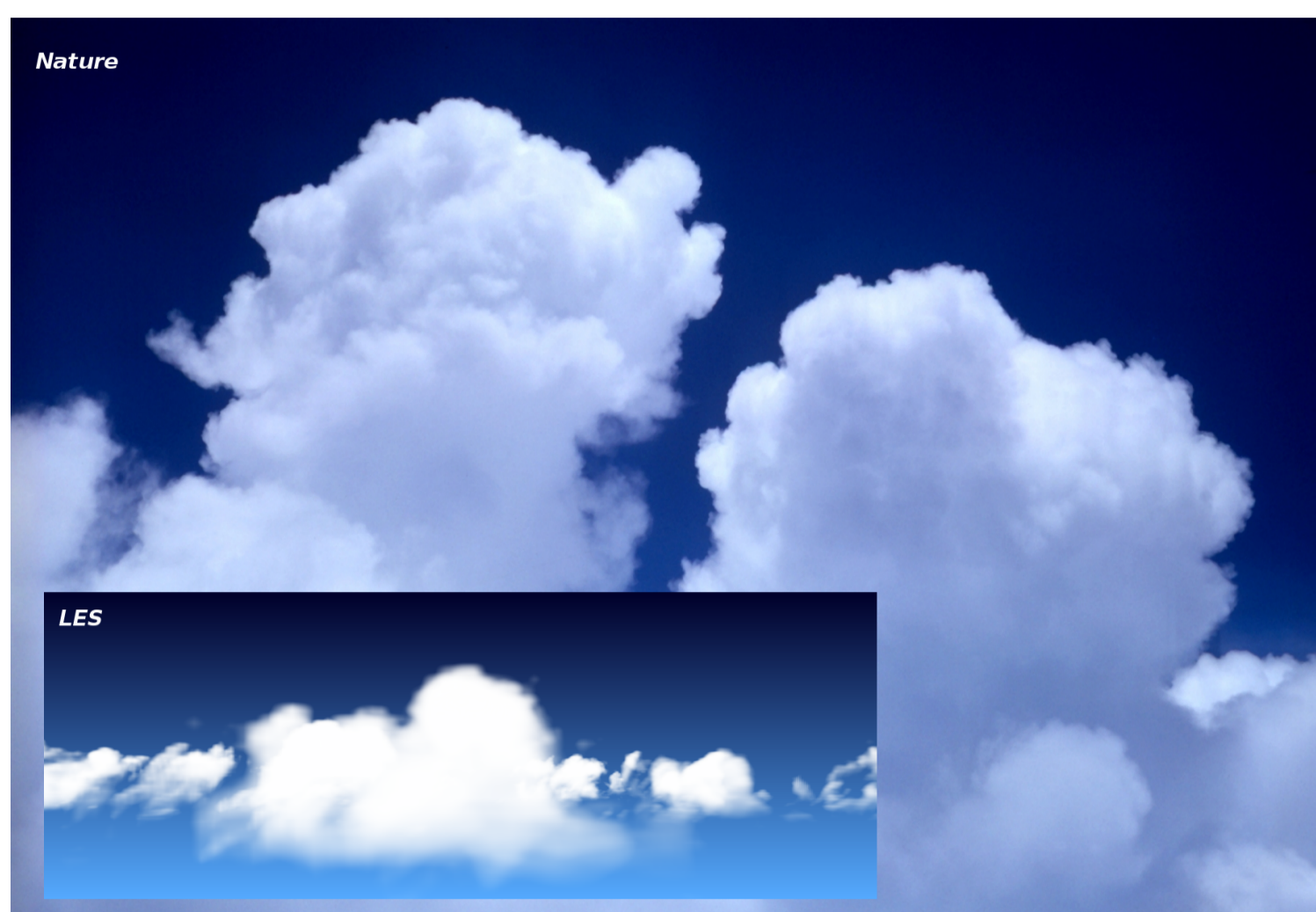


Figure 1 - Real and simulated cumulus clouds.

Large-scale models for weather and climate cannot resolve clouds within a vertical grid column and rely on parameterizations, leading to uncertainty in the representation of clouds and the way they overlap in the vertical. The uncertainty in the cloud overlap remains a significant source of error in the Earth's radiation budget.

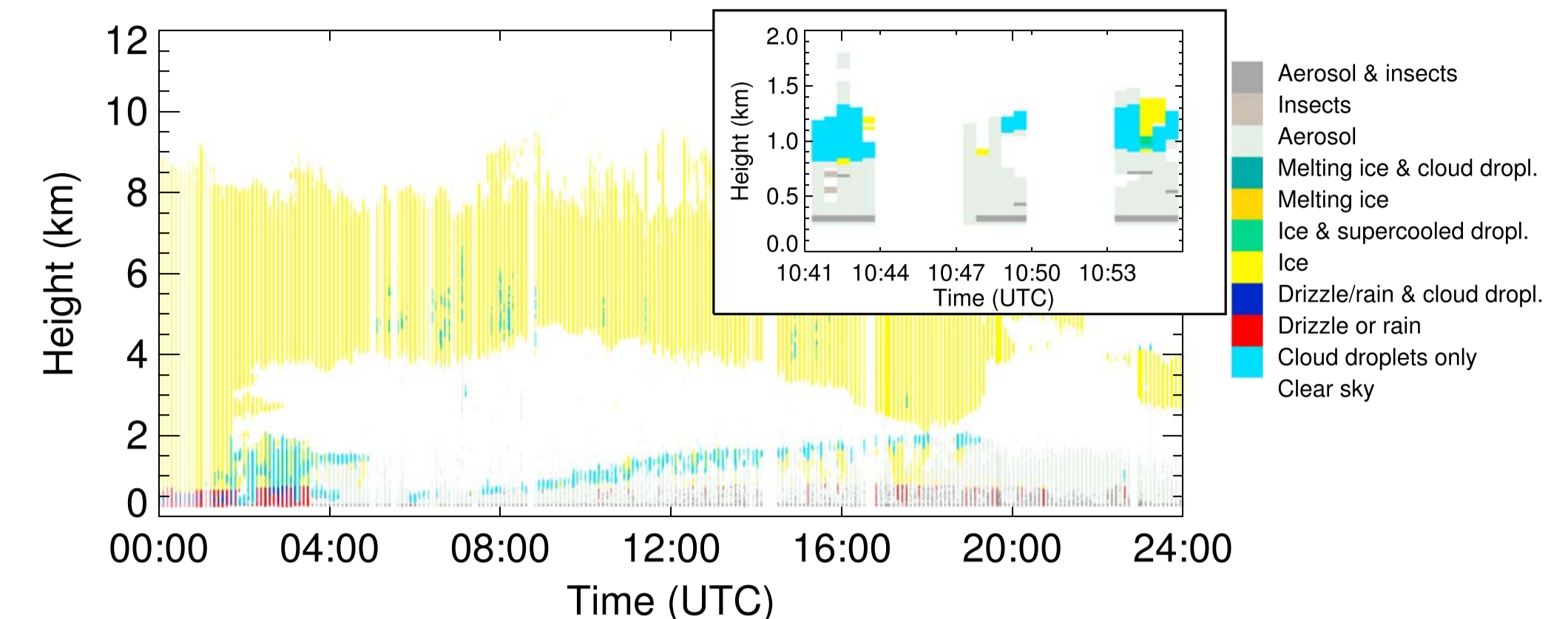
This study aims to compare shallow cumuli vertical overlap statistics derived from high-resolution ground-based measurements and LES simulations.

A better understanding of the unresolved cloud overlap now opens the door for parametrizations, leading to a more reliable cloud radiative budget in large-scale models.

Model and measurements

5 cases featuring boundary layer clouds at JOYCE

Figure 2 - Time-height section of Cloudnet cloud classification on 27 April 2013 at the Joyce supersite. The inset illustrates a model grid with 15 min temporal and 1500 m vertical resolution. Ice pixels are yellow; liquid pixels are light blue.



Dutch Atmospheric LES model

- Homogeneously driven by ECMWF analyses
- Horizontal and vertical resolution: 50 m, 40 m

Cloudnet cloud classification

- "cloud droplets only" pixels are used
- Time and vertical resolution: 30 s, 30 m

Overlap efficiency

To mimic the discretization of large-scale models, daily time-height sections of the cloud masks are divided into equally sized grid boxes, using a temporal resolution of 3 or 15 min (~ 2 and 9 km assuming a wind speed of 10 m/s).

- Cloud cover by area C_a and by volume C_v are calculated using a vertical discretization ranging from 60 to 1500 m for LES and observations.

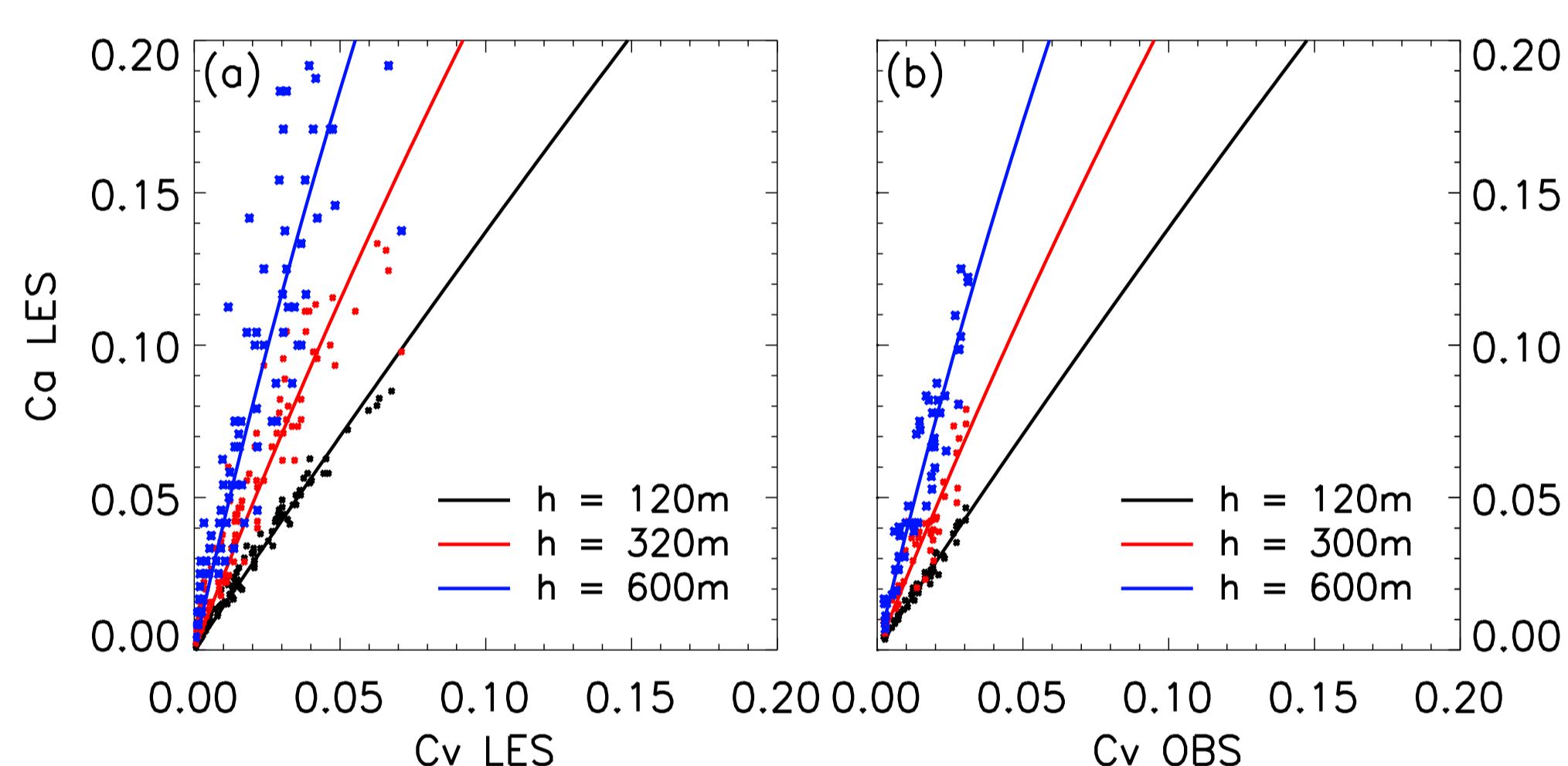


Figure 3 - Scatterplots of C_a vs C_v for 15 minutes time resolution from observations (a) and LES simulations (b) for three different vertical resolutions (120, 320 and 600m).

Cloud overlap efficiency as a function of layer depth h is described by:

1. Overlap ratio R
fit parameter β
 $R = C_v / C_a$
 $R(h) = (1 + \beta h)^{-1}$
2. Overlap parameter α
decorrelation length Z_0
 $C_a = \alpha C_{maximum} + (1 - \alpha)C_{random}$
 $\alpha(h) = \exp(-h / Z_0)$

Inefficient cloud overlap at small scales is supported by observations

- Agreement is found between R derived from observations and simulations.
- Decorrelation lengths are much smaller (< 300 m) than previously reported (> 1 km).
- Including stratiform ice cloud (fig. 2) β drops by one order of magnitude.

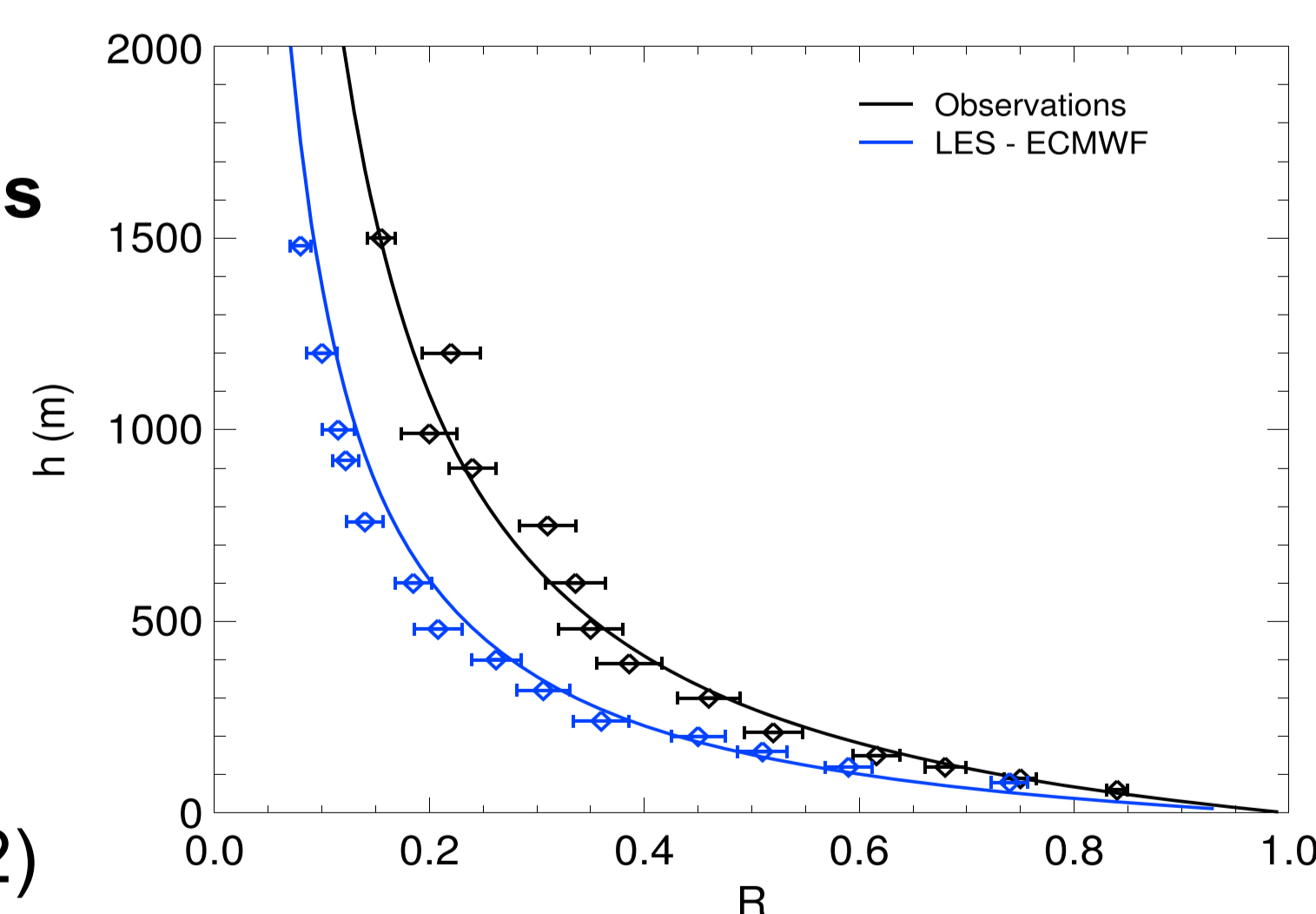


Figure 4 - Daily mean values (diamonds) and inverse linear fitting curves (thick lines) of the overlap ratio R as a function of vertical grid resolution h on 5 June 2013. Error bars are the standard deviation of the mean overlap ratio within the time period considered.

Diurnal cycle

Investigation of the time evolution of β during the diurnal development of the shallow cumulus-capped boundary layer is performed for one case study. The period featuring boundary layer cloud is discretized into 1 h bins, for each of which the overlap ratio values are averaged using a temporal resolution of 15 min.

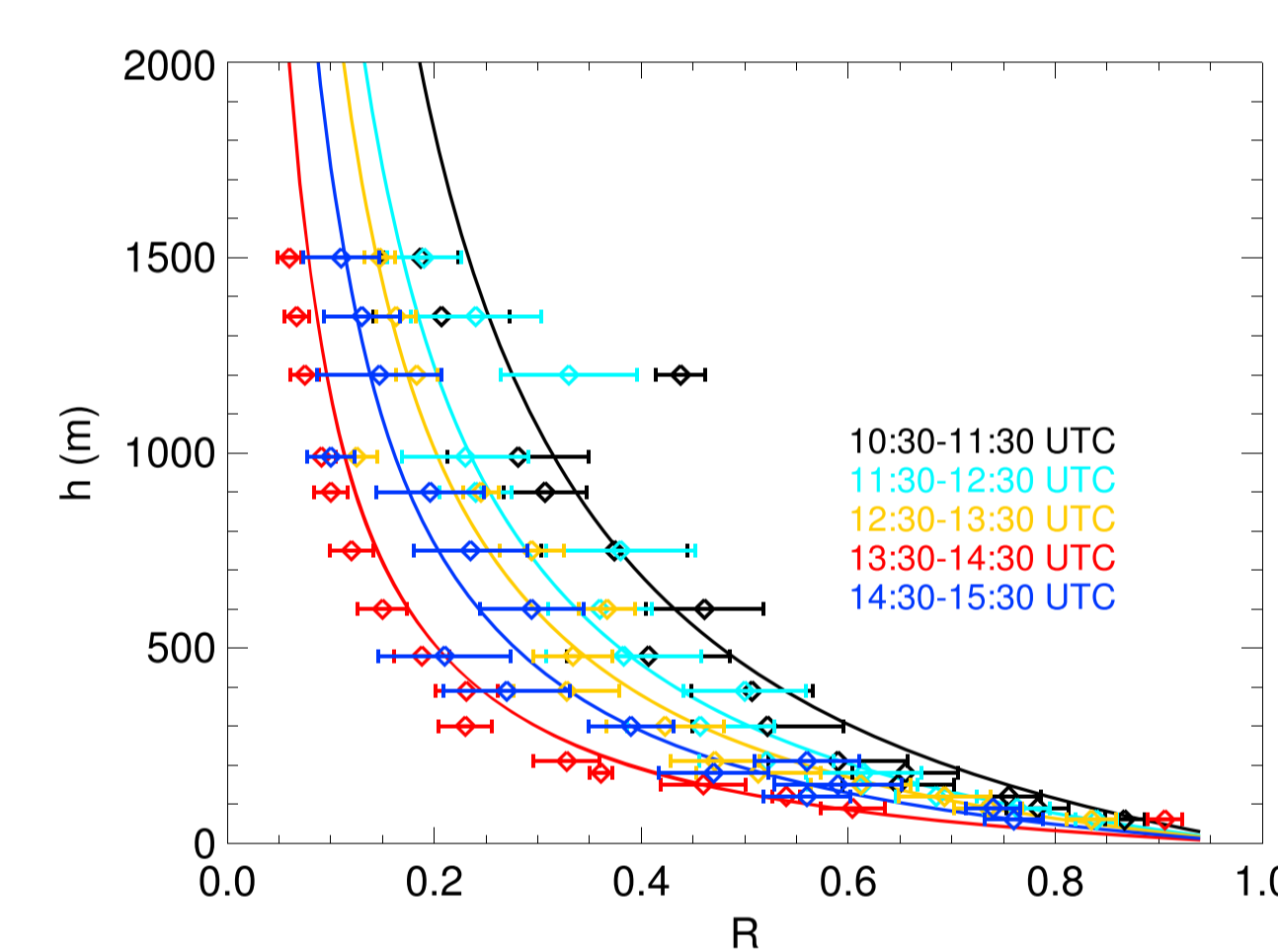


Figure 5 - Daily mean values (diamonds) and inverse linear fitting curves (thick lines) of the overlap ratio R as a function of vertical grid resolution h on 5 June. Error bars are the standard deviation of the mean overlap ratio within the time period considered.

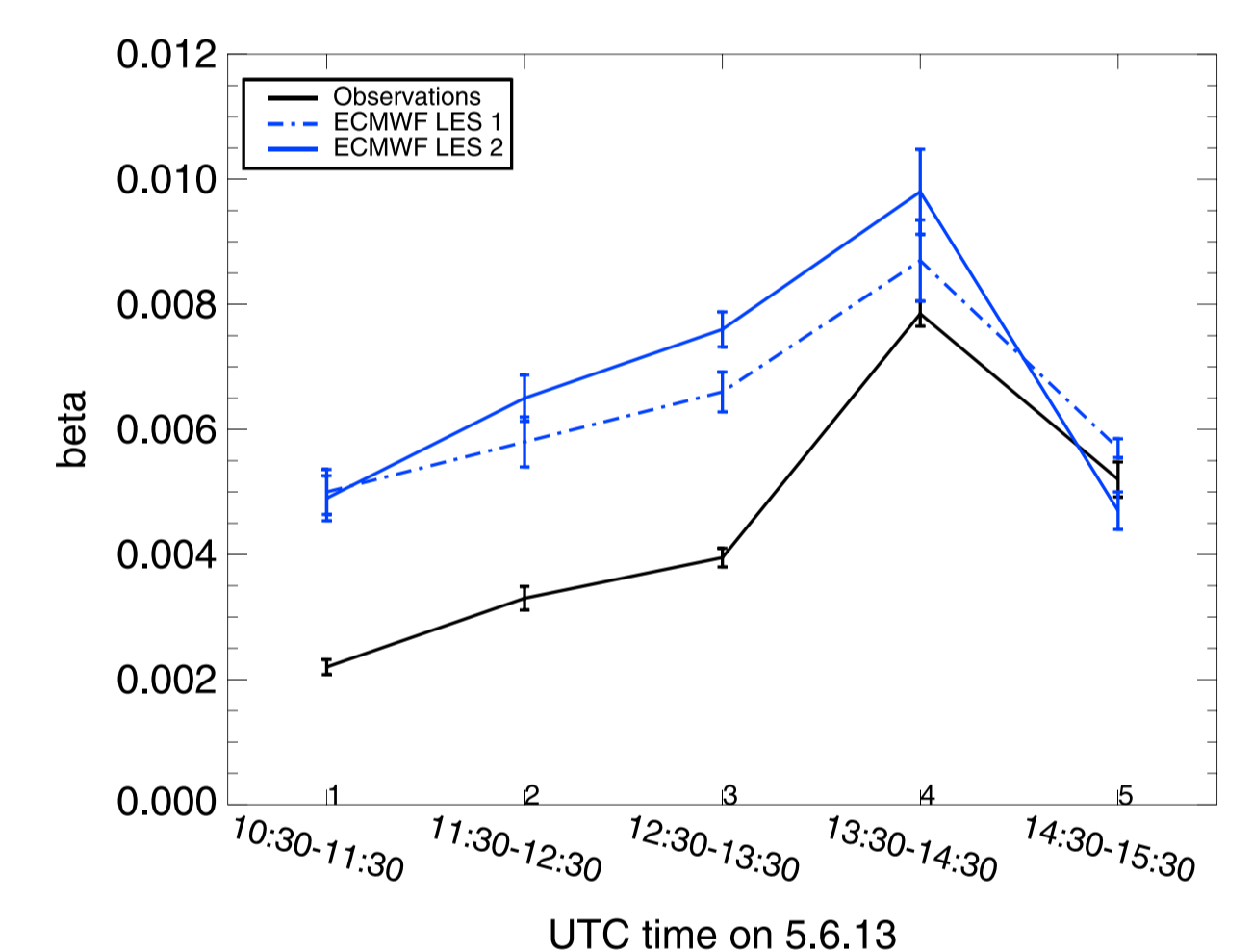


Figure 6 - Time series of fit parameter β (m^{-1}) from observations (black) and ECMWF-driven (blue) LES simulation for two randomly chosen locations. Error bars are standard error of mean β parameter over the time period considered.

- Overlap efficiency, together with cloud cover, decreases with time after cloud onset.
- LES and observations show similar behaviour.
- Previous study reports an opposite behaviour for a case of transient continental cumulus at the ARM SGP site.
- The factor controlling the overlap efficiency evolution remains unclear and will be investigated in future studies.

Conclusion and perspective

- Inefficient overlap for shallow boundary layer clouds is supported by observations and LES.
- Observations and LES depict a decreasing overlap efficiency with time after cloud onset.
- Agreement between observations and LES models suggest their use as a virtual laboratory for parameterization development for larger scale models.



Next step: long-term rigorous verification of LES results at fixed meteorological supersites with continuously operated instrumentation.

Table 1. β Parameter (m^{-1}) (Fitted Using Daily Mean R Values) and Decorrelation Length (m) on Different Days in 2013 Featuring Boundary Layer Clouds Calculated From Observations and LES Simulations for 3 and 15 min Time Resolutions

Day	3 min				15 min			
	$\beta \times 10^3$		Decorrelation Length		$\beta \times 10^3$		Decorrelation Length	
	Observations	LES	Observations	LES	Observations	LES	Observations	LES
27 Apr	4.9	4.5	590	180	-	-	-	-
19 May	5.8	6.2	157	127	-	-	-	-
5 Jun	4.7	6.5	160	148	5.2	6.3	170	202
10 Jun	4.4	4.9	253	104	4.7	5.7	213	153
20 Aug	5.3	6.6	249	120	5.0	7.0	237	239